

COMMENTARY

Can Nitrogen Sequestration Explain the Unexpected Nitrate Decline in New Hampshire Streams?

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Goodale and others (2003) pose an intriguing question: How can we explain the unexpected decline in nitrate concentration in New Hampshire's streams? Theory of forest nitrogen (N) biogeochemistry suggests that retention of atmospherically deposited N can occur through vegetation uptake, accumulation in soil organic matter, abiotic immobilization, and accumulation in coarse woody debris including roots and stumps. The latter mechanism occurs through microbial processes that narrow the C:N ratios of this pool and may be underappreciated, but is well documented (Harmon and others 1986; Fahey and others 1988; Fahey and Arthur 1994). The relative strength of these sinks depends upon site conditions and, in particular, on historical land use (Dise and others 1998; Gundersen and others 1998a; Goodale and others 2000). In general, the greater the previous disturbance and resulting soil N depletion, the greater the soil N sequestration potential. Forest N sinks are limited, thus, N sequestration will not continue indefinitely (Stoddard 1994; Aber and others 2003), but the capacity of this N sink and its dynamics over time are uncertain. As an example of the uncertainty in the temporal dynamics of this N sink, consider the lag between N additions and measurable effects on process. There can be a significant delay between the initiation of inorganic N additions and the induction of detectable increases in net nitrification (Magill and others 2000).

The unexpected decline in nitrate concentrations in New Hampshire streams cannot be explained by trends in forest maturation or atmospheric deposition of N (Goodale and others 2003). Likewise, mechanisms for the decline, based on a re-equilibration following recovery from insect defoliation and severe drought in the 1960s (Aber and others 2002), and soil frost dynamics, were provisionally ruled out based on their failure to explain the regional nature of the stream water nitrate decline (Goodale and others 2003). Potential explanation based on enhanced N uptake through growth increases caused by CO₂ fertilization (Idso 1999) was rejected as having too small an effect on forest growth and the lack of a growth response at the Hubbard Brook Experimental Forest (HBEF). Forest growth data from HBEF (Goodale and others 2003) and Forest Inventory Analysis (FIA) survey data for New Hampshire (FIA data files online) do not support substantial growth increases in recent decades. Goodale and others (2003) concluded that interannual climate variation remained a plausible mechanism; however, they noted that the specific climate factors were uncertain. This region has experienced warmer (Keim and others 2003), wetter (Groisman and others 2001), and longer growing seasons (Cooter and LeDuc 1995; Hodgkins and others 2003) in recent decades, all of which would argue for increased, rather than decreased, N uptake.

There may be another more likely explanation that the authors have not considered. Immobilization of inorganic N (accumulation in organically bound soil pools) in combination with a regionwide

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recovery from past disturbances may explain the decrease in stream water nitrate in recent decades. Net immobilization can occur through several mechanisms: (1) decomposition and microbial transformations of organically bound N in plant residues, (2) microbial uptake of inorganic N and assimilation into microbial biomass that is subsequently transformed into soil organic matter, (3) abiotic N immobilization, (4) microbial transformation of extant soil organic matter that leads to a decrease in C:N ratio, and (5) a suppression of N mineralization. A recent synthesis provides some insight into whether the latter two mechanisms are likely contributing factors. Aber and others (2003) report that soil data from a broad spectrum of northeastern U.S. forests "show an increase in nitrification with decreasing soil C:N ratio." They found weaker relations between N deposition and soil C:N ratio or nitrification. Studies along N-deposition gradients indicate that soil C:N ratios decrease under conditions of chronically elevated N deposition (McNulty and others 1991; Gundersen and others 1998b). These studies are consistent with a model of elevated N deposition resulting in increases in nitrification and decreases in C:N ratio, but they do not support the suppression of N mineralization as a likely mechanism for increasing N sequestration.

Forests across most of New Hampshire, like the HBEF, represent ecosystems that are recovering from major historical disturbances such as logging, grazing, and sometimes cultivation (Cogbill and other 2002; Sundquist and Stevens 1999). Aggrading forests usually have a large capacity to accumulate organic matter as they gradually approach predisturbance equilibrium (Johnson 1992; Huntington 1995). These soil pools include the mineral soil, organic-rich forest floor, and coarse woody debris. Disturbance associated with timber harvesting redistributes high C:N ratio forest floor and coarse woody debris under mineral soil (Ryan and others 1992) that can prime the soil for N sequestration during forest regrowth through microbial transformations. The amount of annual soil N accretion necessary to offset the observed average 71% decrease in stream-water nitrate loss (Goodale and others 2003) (nominally, from ~ 1.2 to $0.3 \text{ g N m}^{-2} \text{ y}^{-1}$ based on nitrate losses at HBEF during the late 1990s) is only about 0.1% of the estimated soil N content at Hubbard Brook (760 g N m^{-2}) (Huntington and others 1990). It seems plausible that forest soils could sustain the accumulation of N at these relatively low levels, which would be very hard to detect upon remeasurement.

The regional soil N sink strength could be greater in recent decades because environmental condi-

tions are now more favorable for soil N sequestration, if not forest growth. Elevated atmospheric CO_2 concentrations intensify carbon cycling that supplies more labile carbon to soil microbes (Hungate and others 1997), increasing capacity for immobilization of inorganic N. Chronically elevated N deposition can intensify N cycling (Galloway and others 2003) that, in turn, can increase net N sequestration rate, provided the supply of labile carbon is not limited. Stimulation of forest growth by CO_2 fertilization, N deposition, and more favorable climate may have exerted more influence on the production of organic carbon belowground (fine root turnover and root exudate) and leaf litter than on stem biomass. Such a differential effect of elevated CO_2 has been demonstrated in Free Air Carbon dioxide Enrichment (FACE) studies (Norby and others 2002). These rapidly cycling pools can contribute to the accumulation of soil organic N (Hungate and others 1997). This proportionately greater stimulation of belowground carbon production could account for an increase in soil organic-N storage without concomitant increase in forest growth. However, it remains uncertain why forest stemwood biomass would not have responded to more favorable climate, CO_2 enrichment, and elevated N deposition in recent decades. Other factors such as growth inhibition due to ozone (Gregg and others 2003) and other stressors or nutrient depletion (Federer and others 1989) may have suppressed stem growth. Abiotic N immobilization is another mechanism of N retention in forest soils (Davidson and others 2003) but its quantitative importance is not yet established. The rate of abiotic N immobilization could be increasing because it has been shown that the ratio of abiotic to total N immobilization is positively related to N concentration and ecosystem N saturation (Johnson and others 2000).

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